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Melting of math xmlns="http://www.w3.org/1998/Math/MathML">mrow>mi >MgSi/mi>msub>mi mathvariant="normal">O/mi> mn>3/mn>/msub>/mrow>/math> determined by machine learning potentials Jie Deng, Haiyang Niu, Junwei Hu, Mingyi Chen, and Lars Stixrude Phys. Rev. B **107**, 064103 — Published 13 February 2023

DOI: 10.1103/PhysRevB.107.064103

### Melting of MgSiO<sub>3</sub> determined by machine learning potentials

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- 11 Keywords:
- 12 Melting, MgSiO<sub>3</sub>, Machine learning
- 13

# 16 Abstract

17 Melting in the deep rocky portions of planets is important for understanding the thermal evolution

18 of these bodies and the possible generation of magnetic fields in their underlying metallic cores.

But the melting temperature of silicates is poorly constrained at the pressures expected in super-Earth exoplanets, the most abundant type of planets in the galaxy. Here, we propose an iterative

21 learning scheme that combines enhanced sampling, feature selection, and deep learning, and

22 develop a unified machine learning potential of *ab initio* quality valid over a wide pressure-

temperature range to determine the melting temperature of MgSiO<sub>3</sub>. The melting temperature of

the high-pressure, post-perovskite phase, important for super-Earths, increases more rapidly with

increasing pressure than that of the lower pressure perovskite phase, stable at the base of Earth'smantle. The volume of the liquid closely approaches that of the solid phases at the highest pressure

27 of our study. Our computed triple point constrains the Clapeyron slope of the perovskite to post-

28 perovskite transition, which we compare with observations of seismic reflectivity at the base of

- 29 Earth's mantle to calibrate Earth's core heat flux.
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# I. INTRODUCTION

37 Cosmic abundances, stellar spectroscopy, observations of polluted white dwarfs, and mass-radius relations point towards the abundance of planets in our galaxy with Earth-like compositions, with 38 39 a mantle dominated by the MgSiO<sub>3</sub> component (~70% in the case of Earth) and an iron-rich core, and masses similar to or greater than that of Earth (1-10 Earth masses) [1,2]. Studies of planetary 40 41 accretion and thermal evolution suggest that these bodies may have begun in a completely molten 42 state and that mantle and core are still partially molten after billions of years [3-6]. The melting 43 temperature of MgSiO<sub>3</sub> exerts a first-order control on thermal evolution because of the large 44 change in viscosity across the melting transition, which sets the time scale for thermal evolution. 45 The density contrast between liquid and solid is also important because this determines whether 46 crystals freezing out of a deep molten portion of a planet sink or float, setting the vector of chemical

47 evolution [7].

48 The melting temperature of  $MgSiO_3$  has therefore attracted considerable attention, yet no consensus exists, in part because of the experimental challenges at very high pressure [8-11]. 49 50 While melting of the bridgmanite phase has received the most attention, melting of its high-51 pressure polymorph – post-perovskite – is also important because this is the stable crystalline phase 52 at pressures greater than 140 GPa (nearly coinciding with the pressure at the base of Earth's mantle) 53 to pressure as high as 750 GPa [12] (the pressure at the base of the mantle in a 5 mass super-54 Earth). The melting temperature also constrains the triple point at which all three phases are stable (bridgmanite, post-perovskite, liquid) and therefore the Clapeyron slope of the solid-solid 55 56 transition, which is observed via seismic reflection at the base of Earth's mantle [13]. The 57 Clapeyron slope of the solid-solid transition is also very uncertain at present, leading to large uncertainties in the heat flux from Earth's core [14], the existence of an active dynamo to generate 58 59 the magnetic field, and the habitability of planets.

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61 Simulation of the MgSiO<sub>3</sub> system at deep Earth and super-Earth conditions presents many 62 challenges. Among these are the subtle energetics of structurally similar phases and the nature of 63 the bonding, which is dominantly ionic, but may also include covalent and metallic contributions, 64 which may account for the discrepancy of previous attempts to simulate melting in this system on the basis of empirical potentials [15-17]. The situation points towards density functional theory 65 (DFT) as an accurate means of representing the energetics of this system that makes no a priori 66 67 assumption about the nature of bonding. However, DFT is very costly, and melting is a rare event in standard molecular dynamics simulations, which is why there have been no ab initio 68 69 determinations of the melting temperature. The solid-liquid two-phase coexistence simulation has 70 been shown to yield robust results for many simpler materials [18-20]. However, the two-phase 71 method requires a large system size to stabilize the coexistence, and very long runs, thus rendering 72 this method computationally demanding or even impossible in the context of the density functional theory.

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75 Machine learning potentials (MLPs) are an emerging approach to atomistic simulations that 76 combines, in principle, *ab initio* accuracy, with the ability to simulate large systems for a very long 77 time [21,22]. Therefore, MLPs are well suited to perform two-phase coexistence simulation and 78 study melting behavior. However, developing a robust MLP of a three-component system of 79 multiple phases over a wide range of pressure and temperature is very challenging [23] and has 80 not yet been attempted. Machine learning potentials are generally trained on DFT results for a finite set of configurations, for example from an *ab initio* molecular dynamics simulation, but there 81 82 are three difficulties. First, the range of thermodynamic conditions sampled by a molecular 83 dynamics simulation is narrow, necessitating multiple simulations covering the pressure-84 temperature range of interest. Second, in any one of these simulations, only one phase will be sampled, since phase transitions are rare events, biasing the MLP towards that phase. Third, 85 86 transitions states, crucial for capturing the physics of the solid-liquid interface are transient and 87 therefore rarely encountered.

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We have overcome these challenges by combining enhanced sampling of configuration space [24] with the deep learning algorithm [22]. The enhanced sampling is driven by a carefully designed set of collective variables (CVs) to capture configurations corresponding to multiple phases, twophase interfaces, and rare transition states. These methods have previously been used to study phase transitions in simpler systems over much narrower ranges of pressure and temperature [23,25]. We develop a unified machine learning potential that encompasses the physics of the three phases of interest over a wide range of pressure and temperature.

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# **II. BUILDING THE MACHINE LEARNING POTENTIAL**

99 A machine learning potential is a non-parametric representation that approximates the Born-100 Oppenheimer potential energy surface to arbitrary accuracy. In our approach, the machine learning 101 potential is trained on a set of configurations drawn from multithermal and multibaric (MTMP) 102 simulations [24], which are used to efficiently sample multi-phase configuration space. The key to driving the sampling is the design of an appropriate collective variable (CV) that captures key 103 104 aspects of the structure. We use an iterative learning scheme to efficiently select distinct samples 105 from the molecular dynamics trajectories. We have found that efficient training of the machine 106 learning potential is greatly facilitated by performing the underlying ab initio calculations at very 107 high precision.

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# A. Multithermal–multibaric simulation

110 111 The multithermal and multibaric simulation is an enhanced sampling technique designed to sample 112 uniformly in energy and volume simultaneously by taking the intervals of temperature and pressure 113 as inputs. It is based on variationally enhanced sampling, where a functional of the bias potential 114 V(s) is introduced as

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$$\Omega[V(\mathbf{s})] = \frac{1}{\beta} \log \frac{\int d\mathbf{s} e^{-\beta[F(\mathbf{s}) + V(\mathbf{s})]}}{\int d\mathbf{s} e^{-\beta F(\mathbf{s})}} + \int d\mathbf{s} p(\mathbf{s}) V(\mathbf{s}) \quad (1)$$

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117 where **s** is a set of collective variables (CV) that are a function of the atomic coordinates **R**;  $\beta = 1/k_B T$  is the inverse temperature with  $k_B$  and T the Boltzmann constant and temperature, 119 respectively;  $F(\mathbf{s})$  is Helmholtz free energy with  $F(\mathbf{s}) = -1/\beta \log \int \delta[\mathbf{s} - \mathbf{s}(\mathbf{R})] d\mathbf{R} e^{-\beta U(\mathbf{R})}$  where  $U(\mathbf{R})$  is the interatomic potential; and  $p(\mathbf{s})$  is a preassigned target distribution. This functional  $\Omega[V(\mathbf{s})]$  is guaranteed to be convex and has a stationary point at

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$$V(\mathbf{s}) = -\frac{\ln p(\mathbf{s})}{\beta} - F(\mathbf{s}) \qquad (2)$$

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124 In this way, one transforms the problem of modifying the Hamiltonian to an optimization problem 125 given the target distribution  $p(\mathbf{s})$ . To generate a multithermal-multibaric ensemble at the pressure 126 and temperature intervals of P<sub>1</sub><P<P<sub>2</sub> and  $\beta_1 > \beta > \beta_2$ , respectively, one chooses the potential 127 energy  $E = U(\mathbf{R})$  and the volume V as CVs to perform a variationally enhanced sampling 128 simulation with the following target distribution

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$$p(E,V) = \begin{cases} \frac{1}{\Omega_{E,V}}, & \text{if there is at least one } \beta \text{ and } P \text{ such that } \beta F_{\beta,P(E,V)} < \varepsilon \\ 0, & \text{otherwise} \end{cases}$$
(3)

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131 where  $\epsilon/\beta$  is a predefined energy threshold. The value of  $\epsilon/\beta$  is set according to the 132 nucleation/melting energy barrier between the solid and liquid states. In practice, we have 133 performed VES simulation using  $s_x$  as CV to roughly estimate the energy barrier, and  $\epsilon/\beta$  should 134 be larger than the energy barrier.

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136 We used PLUMED 2 [26] with variationally enhanced sampling module and LAMMPS [27,28] 137 to perform the multithermal–multibaric simulation on systems of MgSiO<sub>3</sub> consisting of 160 atoms. 138 The bias potential was constructed using variationally enhanced sampling with the energy E, the 139 volume V and  $s_x$  (Eq. 4) as CVs. The basis sets of the bias potential are Legendre polynomials of order 8 for each CV. As a result, there are 729 variational coefficients to be optimized. The 140 141 integrals of the target distribution were performed on a grid of size  $40 \times 40 \times 40$ . Multiple MTMP 142 runs with a pressure interval of 20 GPa are performed. For instance, in the temperature range of 3500 to 5000 K and pressure range of 40 to 60 GPa, the intervals where the polynomials were 143 defined are -108000 < E < -9500 kJ/mol, 1120 < V < 1480 Å<sup>3</sup>,  $-200 < s_x < 3000$ , and the exploration 144 threshold  $\epsilon/\beta$  is set to 150 k<sub>B</sub>T. To improve computing performance, the target distribution was 145 discretized on a grid of dimensions 40×40×40 and smoothed using Gaussians with  $\sigma_E = 200$ 146 kJ/mol,  $\sigma_V = 0.05 \text{ nm}^3$ , and  $\sigma_{sx} = 2$ . The coefficients of the bias potential were optimized every 147 500 steps using the averaged stochastic gradient descent algorithm with a step size of  $\mu = 10$ 148 149 kJ/mol.

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### **B.** Collective variables

For the collective variable (CV), we use the structure factor, which was shown to be an effective CV to drive the first-order phase transition in simpler systems [29]. In our more complex system, we found it essential that the CV contain information from partial structure factors at multiple scattering vectors in order to effectively differentiate bridgmanite, post-perovskite, and liquid as summarized in Fig. 1.



Figure 1. Simulated structure factors (i.e., diffraction intensity) of bridgmanite (a) and liquid. (a) Schematic 161 illustration of the crystal structure of bridgmanite, where the orthorhombic unit-cell is indicated by solid 162 163 box. (b-c) Projected crystal structures of bridgmanite in the x-y and x-z planes, respectively. (d) and (e) 164 display the simulated structure factors of bridgmanite and liquid from the three-dimensional and twodimensional perspectives, respectively. The components of the collective variable, in which the descriptors, 165 166 components of the collective variable  $s_x$ , is highlighted with black arrows. The corresponding planes of the two-dimensional descriptors are denoted in (b-c). The subscript i (e.g., Mg, Si) indicates that only element 167 168 *i* is taken into accounts in the calculations.

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170 The collective variable  $s_x$  to drive the phase transition between bridgmanite and liquid is a linear 171 combination of seven descriptors as

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$$s_x = s_{111_{Mg}}^{3D} + s_{111_{Si}}^{3D} + s_{110_{Mg}}^{xy} + s_{210_{Mg}}^{xy} + s_{002_{Mg}}^{xz} + s_{110_{Si}}^{xy} + s_{210_{Si}}^{xy} + s_{002_{Si}}^{xz}, \quad (4)$$

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174 in which the Debye form of structure factor is employed to calculate each component, i.e.,  $s_{hkl}^{3D}$ , 175  $s_{hkl}^{xy}$ , and  $s_{hkl}^{xz}$  are defined below. Due to the complexity of the system, the contribution of Mg and 176 Si atoms from the three-dimensional (3D) and two-dimensional (2D) in the *x*-*y* and *x*-*z* planes are 177 counted separately (Fig. 1) following [30]. The CV to drive the phase transition between post-178 perovskite and liquid is constructed following the same procedure. For simplicity,  $s_x$  is rescaled 179 to the range of 0 to 1, in which 1 refers to perfect solid state and 0 refers to disordered state with 180 lowest structure factor intensities. 181 The first two descriptors correspond to the first main peak intensities of the structure factors of Mgand Si atoms, respectively, and are calculated with the Debye scattering function:

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$$s_{hkl}^{3D} = \frac{1}{N} \sum_{i=1}^{N} \sum_{i=1}^{N} f_i(q) f_j(q) \frac{\sin(Q \cdot R_{ij})}{Q \cdot R_{ij}} w(R_{ij}),$$
(5)

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in which *q* is the scattering vector, hkl refers to the Miller index of bridgmanite,  $f_i(q)$  and  $f_j(q)$ are the atomic scattering form factors and  $R_{ij}$  is the distance between atoms *i* and *j*. A window function  $w(R_{ij}) = \frac{\sin(Q \cdot R_{ij}/R_c)}{Q \cdot R_{ij}/R_c}$  is used to get a smooth behavior of the structure factor;  $R_c$  (=16 Å) refers to upper limit distance.

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191 The descriptors  $s_{110_{Mg}}^{xy}$ ,  $s_{210_{Mg}}^{xy}$ ,  $s_{210_{Si}}^{xy}$ , correspond to the intensities of the two main peaks 192 of the structure factor of one slice layer which is projected into the *x*-*y* plane, which are given by: 193

$$s_{hkl}^{xy} = \frac{1}{N} \sum_{i=1}^{N} \sum_{i=1}^{N} f_i(q) f_j(q) J_0(Q \cdot R_{ij}^{xy}) w^{xy}(R_{ij}^{xy}) w^z(R_{ij}^z), \tag{6}$$

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in which  $J_0$  is the 0<sup>th</sup> order of the first kind Bessel function,  $R_{ij}^{xy}$  is the distance between atoms *i* and *j* in the *x*-*y* plane;  $w^{xy}(R_{ij}^{xy}) = \frac{1}{1+e^{\sigma(R_{ij}^{xy}-R_c^{xy})}}$  refers to a switching function; and  $R_c^{xy}$  (=10 Å) is the distance cutoff,. In addition, only atoms within a distance cutoff  $R_c^z$  (=3.5 Å) along the z direction are taken into account, and  $w^z(R_{ij}^z) = \frac{1}{1+e^{\sigma(R_{ij}^z-R_c^z)}}$  refers to a switching function to make the descriptor smooth.

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201 Similarly, descriptors  $s_{Mg}^{xz}$  and  $s_{Si}^{xz}$  are calculated as:

$$s_{hkl}^{xz} = \frac{1}{N} \sum_{i=1}^{N} \sum_{i=1}^{N} f_i(q) f_j(q) J_0(Q \cdot R_{ij}^{xz}) w^{xz} (R_{ij}^{xz}) w^y (R_{ij}^y), \tag{7}$$
  
where  $R_c^{xz}$  and  $R_c^y$  are set to 10 and 5.2 Å, respectively.

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In order to validate the sampling effectiveness of such CV (Eq. 4), we have further analyzed the local atomic environment of the configurations with a short-range order parameter. Here we adopted an orientationally targeted order parameters[31] building on the smooth overlap of atomic positions (SOAP). The local environment around an atom is denoted as  $\chi$ , and the associated local density is written as

$$\rho_{\chi}(\mathbf{r}) = \sum_{i \in \chi} e^{\frac{-|\mathbf{r}_i - \mathbf{r}|^2}{2\sigma^2}} \tag{8}$$

in which *i* runs over the neighbors in the environment  $\chi$ ,  $r_i$  are the coordinates of the neighboring atoms relative to the central atom, and  $\sigma^2$  is the variance of the Gaussian functions. Here we set  $\sigma$ to 0.5. In order to measure the difference between the environment  $\chi$  and  $\chi_0$  of the reference

structure that contains *n* reference positions, here the perfect crystal phase is used as the reference

structure. Importantly, the three element Mg, Si and O all have unique local environments. The similarity of two environments are compared by

$$k_{\chi_0}(\chi) = \int d\mathbf{r} \,\rho_{\chi}(\mathbf{r})\rho_{\chi_0}(\mathbf{r}) \tag{9}$$

- A spherical average over all the possible orientations of the referce  $\chi_0$  is then performed to get the
- 216 SOAP kernel. As the orientation of the refere  $\chi_0$  is fixed, the similarity can be trivially performed
- and normalized to

$$\tilde{k}_{\chi_0}(\chi) = \frac{k_{\chi_0}(\chi)}{k_{\chi_0}(\chi_0)} = \frac{1}{N} \sum_{i \in \chi} \sum_{j \in \chi_0} e^{\frac{-|r_i - r_j^0|^2}{4\sigma^2}}$$
(10)

where N is the atom number in the configuration. Such CV is a per atom crystallinity metric of the specific phase considered.

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### C. Iterative learning scheme

224 Due to the vast pressure and temperature range targeted and the complex nature of MgSiO<sub>3</sub> system, 225 we use an iterative training scheme to train and gradually refine the MLP (Fig. 2). Here, one 226 iteration means training a new MLP with updated training set. After seven iterations, we find a sufficiently robust and unified MLP for MgSiO<sub>3</sub> bridgmanite, post-perovskite, and liquid at 0-220 227 228 GPa and 2000-8000 K. We emphasize that the feature selection with PCA and the iterative training 229 is crucial to building a balanced and succinct training set. Indeed, the final training set only consists 230 of 4324 configurations while covering three phases over 6000 K and 220 GPa, compared with 231 typically tens of thousands of frames for mono-atomic species at very narrow pressure and 232 temperature conditions [23,25]. Compared with other active learning algorithm like DP-GEN, the 233 iterative learning scheme presents two improvements: 1) efficient enhanced sampling is embedded 234 in the workflow; 2) we use PCA analysis and comparison between the MLP prediction and VASP 235 results (ground truth) to select the candidate frames, while DP-GEN relies on the model deviation 236 of the candidate frames only. We found that the model deviation, although being computationally 237 more efficient, is prone to selecting frames that are already predicted very well by the MLP, 238 especially when the threshold of the model deviation is not well set.



241 Figure 2. Flowchart of the iterative training scheme. 1) We first build an initial dataset for liquid 242 phase only and train a preliminary MLP as reported in our recent study [32]. Liquid, as a disordered phase, may encompass some of the local environments of the solid phases, and thus may serve as 243 a good starting point for generating a unified MLP for both solid and liquid phases. 2) 244 245 Multithermal-multibaric (MTMP) simulations are performed with the MLP using LAMMPS 246 interfaced with PLUMED 2. Here, PLUMED 2 is used to calculate the CVs and implement the 247 enhanced sampling method. The target pressure and temperature ranges are very large in this study, 248 making it difficult to cover in one multithermal-multibaric simulation. We find that MTMP 249 simulations with ~20 GPa and ~2000 K intervals yield good convergence and can sample the phase 250 transition sufficiently. As a result, the target P/T ranges are divided into 20 GPa and 2000 K bins 251 along the melting curves of bridgmanite and post-perovskite [4,33]. We gradually update the P/T intervals with the iteration. 3) The resulting trajectories are saved every 500 timesteps. The saved 252 253 frames are converted to design matrices based on the smooth overlap of atomic positions descriptor [34]. We then perform principal component analysis on these design matrices and select the 254 255 candidate configurations using the farthest point sampling technique [35,36]. The size of candidate configurations N is large at first a few iterations and gradually decreases at later iterations. 256 257 Specifically, N=1000-2000 for the first two iterations, N=50-100 for the rest of iterations. 4) The 258 selected frames were recalculated with DFT. The resulting energies and forces, are compared with 259 the MLP predicted ones. For the first three iterations, the configurations with both energy 260 difference >15 meV/atom and atomic forces difference >0.5 eV/Å are selected. For the rest of 261 iterations, we relax the selection criteria to energy difference >5 meV/atom and atomic forces difference >0.25 eV/Å. The size of the selected configurations in this step is M. We emphasize the 262 selection criteria here is unlikely to be universal for all other systems but the principle that relaxing 263 264 the selection criteria with iterations should apply. 5) The selected configurations will be combined with the initial dataset to train a new MLP. We re-iterate above steps until we cannot select frames 265 in step (4) (i.e., M=0) and all the target pressures and temperatures are covered by MTMP 266 267 simulations in step 2). 268

### D. DeePMD approach.

271 The DeePMD approach adopts an end-to-end strategy [22,37] and does not make a priori 272 assumptions about the form of the descriptor but rather uses a deep neural network to determine its form based only on the spatial location, in a suitably defined coordinate frame, of the 273 274 neighboring atoms. Neural networks are widely used in the development of machine learning 275 potentials because they are, in principle, capable of approximating any continuous function to 276 arbitrary accuracy [32,38,39]. With DeePMD, one uses a neural network to find the functional 277 form of the descriptor and a second neural network (fitting network) to determine the form of the 278 potential energy surface. The fitting network is composed of three layers with 240 nodes in each 279 layer. A cutoff of 6 Å is employed to describe the atomic local environments. The loss function is 280 defined as

$$L(p_{\epsilon}, p_{f}, p_{\xi}) = p_{\epsilon} \Delta \epsilon^{2} + \frac{p_{f}}{3N} \sum_{i} |\Delta F_{i}| + \frac{p_{\xi}}{9} ||\Delta \xi||^{2}, \qquad (11)$$

where  $p_{\epsilon}$ ,  $p_f$ ,  $p_{\xi}$  are tunable prefactors for difference between the MLP prediction and training data. *\epsilon* is the energy per atom;  $F_i$  atomic force of atom *i*;  $\xi$  the virial tensor divided by N; N the number of atoms. We adopt the conventional setting of increasing both  $p_{\epsilon}$  and  $p_{\xi}$  from 0.02 to 1 while decreasing  $p_f$  from 1000 to 1 over the course of training.

E. Ab initio calculations

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289 All *ab initio* calculations were performed on MgSiO<sub>3</sub> consisting of 160 atoms based on DFT in the 290 PBEsol approximation [40] using VASP [41]. We used the projector augmented wave (PAW) 291 method [42] as implemented in VASP [41]. We use the PBEsol approximation as it has been found 292 to yield good agreement with experimental measurements of physical properties of silicates and oxides [43-45] and melting temperatures of MgO [46]. The core radii are O: 0.820 Å (2s<sup>2</sup>2p<sup>4</sup>), Si: 293 294 1.312 Å (3s<sup>2</sup>3p<sup>2</sup>), Mg: 1.058 Å (2p<sup>6</sup>3s<sup>2</sup>). To construct the initial dataset, we perform *ab initio* molecular dynamics (AIMD) simulations with relatively low precision settings: energy cutoff of 295 500 eV, energy cutoff of 10<sup>-4</sup> eV, and Gamma-point only k-mesh. AIMD simulations are 296 297 performed in the NVT ensemble (constant number of atoms, volume, and temperature) using the 298 Nosé-Hoover thermostat [47] and run for 5-20 ps with 1 fs time step. We assume thermal 299 equilibrium between ions and electrons via the Mermin functional [48].

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301 The configurations generated by these AIMD simulations as well as the multithermal and/or 302 multibaric MD simulations were then selected to construct the MLP. The energy, force, and stress of these selected configurations were recalculated at much higher precision with: the energy cutoff 303 304 that sets the size of the basis set increased from 500 eV to 800 eV, the precision to which the selfconsistent solution to the Kohn-Sham equations is found lowered form 10<sup>-4</sup> eV to 10<sup>-6</sup> eV, and 305 306 sampling of the Brillouin zone increased from the Gamma-point only to a 2×2×2 Monkhorst-Pack mesh. We found this high precision recalculation to be essential for optimizing the accuracy and 307 308 scope of the MLP [32] (Fig. 3).



Figure 3. Convergence tests of total energy (a) and pressure (b) with varying energy cutoff
(ENCUT flag in VASP) for 32 MgSiO<sub>3</sub> bridgmanite at static condition. An energy cutoff of 800

eV is sufficient to obtain converged results for both energy and pressure.

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# **F. Simulations**

316 For two-phase simulations, we use LAMMPS to build a 2-phase model of coexisting solid and 317 liquid with the ratio of 1:1. Supercells of 900 atoms are constructed ( $3 \times 3 \times 5$  for *Pbnm* bridgmanite and  $3 \times 5 \times 3$  for *Cmcm* post-perovskite) and then relaxed for 1000 steps at the desired pressure 318 319 and temperature conditions in the NPT ensemble. The relaxed cell is then used to perform NVT 320 simulations at high temperatures far exceeding the melting temperatures with the atoms of half the cell fixed and the force applied to these atoms set to be 0. The resulting structure is half-molten 321 322 and half-crystalline. We relaxed this structure again at the desired pressure and temperature 323 conditions for 1000 steps to obtain the initial configuration for two-phase simulations. We also 324 tested the size effect on the melting temperature by exploring larger system sizes of 1800 atoms 325 and 3000 atoms, and found that systems of 900 atoms are sufficient to yield identical melting 326 points as those larger systems.

327 Simulations on the two-phase cell were performed at the desired pressure and temperature 328 conditions (NPT). If the whole cell is molten (crystallized) in the end, the simulation temperature 329 is above (below) the melting point. The state of the system can be determined by interrogating the 330 radial distribution functions. In this way, we can obtain the upper and lower bounds of the melting 331 curve.

332Phonon dispersions and zero-point energy were performed using the PHONOPY program [49].333Real-space force constants were calculated with density functional perturbation theory [50], with334 $2 \times 2 \times 2$  and  $4 \times 1 \times 2$  supercells for bridgmanite and post-perovskite, respectively.

- 335 336 **III. RESULTS**
- 337338A. Sampling of Configuration Space

340 Our approach yields a broad sampling of configuration space and an efficient selection of 341 representative configurations (Fig. 4). From a single MTMP run (40-60 GPa, 3000-5000 K), we 342 generate liquid and bridgmanite configurations as well as configurations containing an interface between the two phases, and configurations containing distorted and defective crystalline 343 344 structures. Post-processing with principal component analysis (PCA) selection and an iterative 345 learning scheme yield balanced and succinct sampling over this range (Fig. 4). Indeed, the final 346 training set is very small, consisting of only 4324 configurations, while covering a wide 347 temperature and pressure (2000 to 8000 K and 0 GPa to 220 GPa), compared with typically tens 348 of thousands of frames for mono-atomic species over much narrower ranges of pressure and 349 temperature conditions [23,25].

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352 Figure 4. Configurations explored by multi-thermal multibaric simulations at 40-60 GPa and 353 3000-5000 K. The energy of the system (160 atoms) is plotted against volume and color-coded 354 by the value of collective variable (CV) defined in Eq. 4. Large and small CVs indicate a 355 perovskite-like or liquid-like state, respectively. Snapshots of configurations are shown in the 356 circles with atoms color-coded based on the orientationally targeted order parameters [31] building 357 on the smooth overlap of atomic positions (SOAP) [34], with red indicating perovskite-like and 358 blue indicating liquid-like local atomic environments. The yellow and green ellipses show the 359 regions sampled by standard molecular dynamics simulations at the same P/T range for liquid and 360 solid states, respectively.

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# **B. Benchmarks of the Machine Learning Potential**

We compare the energies, atomic forces, and stresses from the machine learning potential with those from DFT simulations for 35585 configurations that are not included in the training set (Fig. 5). The root-mean-square errors of the energies, atomic forces, and stresses are 4.9 meV/atom, 0.24 eV/Å, and 0.37 GPa, respectively. These uncertainties are comparable to the typical precision of *ab initio* molecular dynamics simulations [45]. As all the testing structures supercells, the robustness of the MLP in predicting properties larger systems is unclear. We performed another
verification test with data obtained using a larger supercell with 320 atoms. This structure is not
included in the training set. The RMSE of energy prediction is similar to the error in the testing
sets (Fig. 6). This verification test further proved the accuracy of energy prediction and also
demonstrated the transferability of the MLP to larger structures.

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**Figure 6.** Comparisons of the total energy changes along molecular dynamics trajectories between the DFT (thick colored lines) and MLP potential (thin black lines) for MgSiO<sub>3</sub> bridgmanite (blue), post-perovskite (red), and liquid (green) at 140 GPa and 5000 K. The models used in this simulation contain 320 atoms, and none of the structures in the trajectories were included in the training set. The root mean square error of MLP is 4.2, 2.4, 7.1 meV/atom for perovskite, postperovskite, and liquid, respectively.

### **C. Two-Phase Simulations**

Two-phase simulations yield precise determinations of the melting temperature. Starting with a two-phase simulation cell of bridgmanite at 140 GPa, a simulation at 6100 K melts completely after 260 ps, whereas a simulation at 6000 K crystallizes after 730 ps (Fig. 7). Post-perovskite, on the hand, melts at 6000 K and crystalizes at 5900 K (Fig. 8). These simulations, performed at constant pressure and temperature show that the system expands upon melting, and contracts upon crystallization, yielding the volume of melting, the Clapeyron slope ( $\Gamma=dP/dT_m$ ) and the entropy of melting ( $\Delta S_m=\Gamma\Delta V_m$ ) (Table 1).



Figure 7. Machine learning molecular dynamics simulations of the coexistence of MgSiO<sub>3</sub>
 bridgmanite and liquid at 140 GPa and 6100 K (upper panel) and 6000 K (lower panel). The
 simulation cell contains 600 MgSiO<sub>3</sub> formula units (3000 atoms). The simulation timestep and the
 corresponding cell shape are also shown.



Figure 8. Machine learning molecular dynamics simulations of the coexistence of MgSiO<sub>3</sub> post perovksite and liquid at 140 GPa and 5900 K (upper panel) and 6000 K (lower panel). The
 simulation cell contains 600 MgSiO<sub>3</sub> formula units (3000 atoms). The simulation timestep and the
 corresponding cell shape are also shown.





**Figure 9.** Melting of MgSiO<sub>3</sub> bridgmanite (a), post-perovskite (b), and the phase boundary 408 409 between bridgmanite and post-perovskite (c). Results from this study are shown solid blue circles for bridgmanite, red circles for post-perovskite, green circles for triple point and zero-K transition 410 point. The uncertainties for melting temperatures are 50 K. The solid colored lines in (a) and (b) 411 represent the Simon fit. The green dashed line in (a), (b), and (c) is the bridgmanite-post-perovskite 412 413 phase transition boundary. Blue, green, and red shadings cover the stability fields of liquid, 414 bridgmanite, and post-perovskite, respectively. (a) Previous experimental results on melting of 415 MgSiO<sub>3</sub> bridgmanite are denoted by upward triangles [8], squares [51], leftward triangles [52], rightward triangles [53]. Experimental results of bridgmanite containing ~10 mol% Fe are shown 416 in solid dark upward triangles [8], downward triangles [9], and stars [54]. Prediction based on 417

418 Lindemann law is shown dotted line [8]. Estimates based on atomistic modeling includes two-419 phase simulations based on classical potential with corrections [15] (dashed line), molecular

420 dynamics simulations with empirical overheating correction [16] (loosely dashed line), and [17]

421 (thin diamond), the integration of Clausius-Clapeyron equation by [33] (dashed-dotted line). (b)

422 Previous results for the melting of post-perovskite include two shock compression experiments

423 [11] (upward triangle) and [10] (downward triangle), the inferred melting curve using the

424 Lindemann law [4] (dotted line), and two-phase simulations based on classical potential with

425 corrections [15] (dashed line). (c) The results of sub-solidus experiments of MgSiO<sub>3</sub> [55,56] are

shown with upward triangles, downward triangles, and squares denoting bridgmanite-only, post-

427 perovskite-only, and bridgmanite-post-perovskite coexistence, respectively.

428

429	Table 1. Calculated	melting propertie	es of MgSiO3: pro	essure <i>P</i> , melting	temperatures $T_1$	m, slope of
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430 the melting curve dT/dP, volume ( $\Delta V_m$ ) and entropy ( $\Delta S_m$ ) of melting at the melting point.

431 Entropy is shown in  $Nk_B$  unit where N is the number of atoms per formula unit and  $k_B$  is the

432 Boltzmann constant.

Phase	P (GPa)	$T_{\rm m}\left({\rm K}\right)$	dT/dP (K/GPa)	$\Delta V_{\rm m}$ (Å <sup>3</sup> /atom)	$\Delta S_{\rm m}(Nk_{\rm B})$
	20	2875(50)	95.1(4.7)	1.26(0.02)	0.95(0.04)
	40	4000(50)	38.3(1.2)	0.71(0.01)	1.33(0.04)
Bridgmanite	75	5000(50)	21.2(0.6)	0.43(0.02)	1.46(0.04)
	120	5750(50)	14.3(0.5)	0.27(0.02)	1.38(0.04)
	140	6050(50)	12.6(0.4)	0.21(0.02)	1.20(0.04)
	160	6250(50)	11.4(0.4)	017(0.01)	1.11(0.04)
	120	5600(50)	17.2(3.1)	0.35(0.01)	1.49(0.27)
	140	5950(50)	15.5(2.7)	0.31(0.01)	1.45(0.25)
Post-perovskite	160	6200(50)	14.2(2.5)	0.27(0.02)	1.41(0.24)
	180	6450(50)	13.1(2.2)	0.25(0.02)	1.37(0.23)
	200	6750(50)	12.2(2.1)	0.23(0.02)	1.33(0.23)

433

434 Our results agree well with a number of experimental studies at the low-pressure end of the 435 bridgmanite stability field (Fig. 9). In order to better constrain the slope of the bridgmanite melting curve, we have performed a simulation at 20 GPa, below the stability field of bridgmanite, but 436 437 accessible to our simulations because of kinetic hindrances to crystal-crystal transitions. At 20 GPa, 438 we obtained a melting temperature of  $2875\pm50$  K, slightly higher than the result (i.e.,  $2700\pm50$  K) of the multi-anvil experiments [51], suggesting that PBEsol may overestimate the melting 439 temperature consistent with the previous study on MgO melting [46]. Our melting curve of 440 MgSiO<sub>3</sub> bridgmanite may be expressed by the Simon equation,  $T_m = 2875 \pm 50 \left( \frac{P-20}{8.11+0.37} + \right)$ 441

442  $1)^{\frac{1}{3.73\pm0.06}}$ , where  $T_{\rm m}$  is in K and P in GPa. This fitted melting curve agrees very well with the 443 results of some laser-heated diamond anvil cell experiments [8,52] up to 50 GPa. The resulting 444 melting slope at 25 GPa is around 69 K/GPa, broadly consistent with experimental result of ~80 445 K/GPa by ref. [8], but deviate from the results of ref. [9,54] (~0 K/GPa), and those of ref. [51] (30

- 446 K/GPa), implying a vanishing small volume of melting, contrary to our findings and that of
- 447 previous *ab initio* determination of the volume of melting [33]. Our results agree well with the
- 448 only determination of the melting temperature at pressure greater than 100 GPa from shock wave
- experiments [53].
- The melting temperature of post-perovskite increases more rapidly with increasing pressure than that of bridgmanite (Fig. 9b); our results can be represented by  $T_m = 5600 \pm 50 \left(\frac{P-120}{113.60\pm 16.13} + \right)$
- 452  $1)^{2.85\pm0.32}$ . Our melting curve is consistent with a shock wave measurement of melting at 500 GPa 453 [10], but is significantly higher than that determined in another study at 210 GPa [11]. We note 454 that extrapolating the melting curve beyond ~200 GPa is subject to uncertainty: for example, our 455 best-fit melting curve of post-perovskite predicts a melting point of 9376±656 K at 500 GPa. More
- 456 experiments at these extreme conditions are clearly warranted.
- 457 From our simulations, we also determine the volume of melting, and from the Clausius-Clapevron
- relation, the entropy of melting (Table 1, Fig. 10). The volume of melting diminishes rapidly with
- 459 increasing pressure because the liquid is more compressible than the solid. The entropy of melting
- 460 initially increases with increasing pressure and then decreases with increasing pressure at pressure
- 461 greater than 100 GPa. The volume and entropy of melting increase at the triple point as the volume
- and entropy of the post-perovskite are less than those of the bridgmanite phase.



464 Figure 10. The volume (top) and the entropy (bottom) of melting from the bridgmanite (blue)465 and the post-perovskite (red) phases. For comparison we also show experimental values for

466 melting from the low-pressure enstatite structure [57,58], and theoretical results for the melting

- 467 of monatomic systems interacting with inverse-power repulsion with the value of the power
- 468 indicated ranging from 1 (one-component plasma) to infinite (hard-spheres)[59].

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# 470

# D. Bridgmanite to post-perovskite transition

The intersection of the melting curves of bridgmanite and post-perovskite yields the 471 bridgmanite/post-perovskite/liquid triple point at 180 GPa and 6420 K (Fig. 9). We combine this 472 473 result with our computed values for the phase transition pressure at 0 K (96 GPa), the volume contrast between the two phases (0.467 Å<sup>3</sup>/formula unit) and the Einstein temperatures of the two 474 phases (773 K for bridgmanite and 791 K for post-perovskite) to determine the solid-solid phase 475 boundary following the formalism of Jeanloz (1989) which accounts for the vanishing Clapeyron 476 slope in the limit of zero temperature (Fig. 9c). The resulting Clapeyron slope at 2000 K is 13.9 477 MPa/K, very close to the experimental result of 13.3±1.0 MPa/K [55], although we note that the 478 479 experimental value may have considerable systematic uncertainty due to the non-unique choice of 480 pressure scale [14].

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# **IV. DISCUSSION**

# A. Melting Curve and Crystal Buoyancy

486 The Lindemann law has been widely used to predict the melting curve of materials including at high pressure [60,61]. We find that for both bridgmanite and post-perovskite, the Lindemann law 487 tends to predict much larger melting slopes  $(dT_m/dP)$ , thus leading to extremely high melting 488 temperatures at high pressure (Fig. 9). The difference between our results and the Lindeman law 489 reveals the importance of liquid structure. Whereas the Lindemann law can be derived by assuming 490 491 that scaled liquid structure is constant along the melting curve. We find, in agreement with previous studies, that liquid structure changes substantially with increasing pressure [33]. 492 493 Moreover, the change in the liquid strucutre, including increases in the Si-O coordination number, 494 are such as to cause the liquid to become denser with increasing pressure, thus decreasing the 495 volume of melting and the Clapeyron slope.

The entropy of melting determined here (Table 1, Fig. 10) is much larger than that of many monatomic systems at high pressure (*R*ln2 where *R* is the gas constant) [62]. We attribute the larger entropy of melting to the range of different Si-O coordination environments in the liquid present at all pressures, producing a liquid structure that is much richer than the nearly closepacked structures of monatomic liquids. At the highest pressures of our simulations, the entropy of melting decreases with increasing pressure, but is still much larger than Rln2, and larger than that of simple monatomic liquids interacting with inverse-power repulsion.

The volume of the liquid closely approaches, but does not fall below that of the solid phases (Table 1, Fig. 10). The volume of melting at the highest pressure of our study (4 %) is much less than that at ambient pressure (18 %), reflecting the greater compressibility of the liquid as compared with solid phases. The very small volume of melting that we find means that crystals freezing out of deep molten portions of rocky planets, are likely to be buoyant. With similar volumes, the liquid is likely to be denser than coexisting crystals because major heavy elements, like Fe, tend to partition favorably into the liquid. Deep crystal buoyancy in cooling rocky planets has important implications for understanding their thermal and chemical evolution [7,63].

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### **B.** Thermal structure of Earth's lowermost mantle

To examine whether the bridgmanite to post-perovskite transition may be encountered in the deep
Earth, we assume that the geotherm consists of an adiabat and a lower thermal boundary layer,
following previous studies [13]

517

$$T(r) = T_{CMB} - (r - r_{CMB})T'_{S} - \Delta T \operatorname{erf}\left(\frac{r - r_{CMB}}{\delta}\right)$$
(12)

518

519 where T(r) is the temperature as a function of radius, subscript *CMB* indicates values at the core-520 mantle boundary,  $\Delta T$  and  $\delta$  are the temperature contrast and thickness of the thermal boundary 521 layer, respectively, and  $T'_{S}$  is the adiabatic gradient at the base of the mantle. The pressure of the 522 CMB is 136 GPa and temperature ( $T_{CMB}$ ) is assumed to be 4000 K. We compute the adiabat  $T_{S}(r)$ 523 for a given potential temperature from HeFESTo [64-66], yielding the adiabatic gradient  $T'_{S}$  and 524 the temperature contrast  $\Delta T = T_{CMB} - T_{S}(r_{CMB})$ . The heat flux at the core-mantle boundary is 525 then

 $F_{CMB} = k \left( T_S' + \frac{2}{\sqrt{\pi}} \frac{\Delta T}{\delta} \right) \tag{13}$ 

526

527 with the thermal conductivity k=8.1 W/m/K[67].

528

529 Our results for the bridgmanite to post-perovskite phase transition suggests that post-perovskite 530 may exist as a lens in the cooler parts of the deep lower mantle (Fig. 9). Along an average mantle 531 geotherm with the potential temperature  $T_P=1600$  K and a bottom thermal boundary layer of greater than  $\delta$ =120 km thickness, only bridgmanite is stable and post-perovskite is absent, 532 533 consistent with the absence of observations of seismic reflections from the lowermost mantle in 534 most regions [68]. If we examine a geotherm representative of a cooler portion of the mantle  $(T_P)$ =1300 K, d=170 km), we find two crossings of the bridgmanite to post-perovskite transition, at 535 536 100 and 300 km above the core-mantle boundary. The depths to these reflections are consistent 537 with paired reflections of opposite polarity seen in cooler parts of the mantle [13,69]. The presence of a post-perovskite lens therefore places important constraints on the thermal structure 538 of the lower-most mantle and on the heat flow emanating from the core. For our cool model 539 540 geotherm, we find a heat flux into the base of the mantle of  $100 \text{ mW/m}^2$ , or 15 TW. This heat flux is more than sufficient to drive dynamo action in the underlying outer core [70]. In detail, the 541 542 pressure-temperature conditions of the bridgmanite to post-perovskite transition may depend on 543 the concentration of Fe, Al, and other secondary oxides [71]. Moreover, the heat flux into the 544 mantle is likely to vary greatly laterally, and may approach zero in some regions of the mantle [67]. Our result for the cool mantle geotherm is therefore consistent with a global average heat flux of 545

 $80 \text{ mW/m}^2$  or 12 TW, which agrees with estimates of the minimum heat flux required to drive magnetic field generation in the underlying core. V. Conclusion It is now possible to develop machine learning potentials that accurately capture the physics of multiple phases in multi-atom systems over a wide range of pressure and temperature. We have overcome the challenges posed by planetary-scale applications with an iterative training scheme that entails multithermal-multibaric enhanced sampling driven by structure factors as collective variables, feature selection, deep learning, and DFT calculations. This scheme allows us to build a MLP of MgSiO<sub>3</sub> liquid, bridgmanite, and post-perovskite up to 220 GPa and 8000 K using only 4324 training frames. The phase stability relations that we have determined using this machine learning potential place important constraints on processes in the deep Earth and in super-Earth exoplanets. Liquids are likely to be denser than coexisting solids within the bridgmanite and post-perovskite pressure-ranges of stability. The Clapeyron slope of the bridgmanite to post-perovskite transition indicates the presence of double-crossings of the phase transition in colder portions of the mantle, consistent with seismic observations and heat flux from the core-mantle boundary compatible with magnetic field generation in Earth's outer core. 

# 570 Data Availability

- 571 The main data supporting the findings of this study are available within the paper and its
- 572 supplementary information. Data set used in this study has been deposited at the Open Science
- 573 Framework (https://osf.io/dt4xs/) with doi: 10.17605/OSF.IO/DT4XS. The software packages
- used in this study are standard codes: VASP (version 5.4) is a commercial code
- 575 (see <u>www.vasp.at</u>) whereas DeePMD-kit (<u>https://github.com/deepmodeling/deepmd-kit</u>),
- 576 PHONOPY (<u>http://phonopy.github.io/phonopy/</u>), LAMMPS (<u>https://www.lammps.org/</u>),
- 577 PLUMED 2 (<u>https://www.plumed.org/doc-v2.6/user-doc/html/index.html</u>), HeFESTo
- 578 (https://github.com/stixrude/HeFESToRepository) and ASAP
- 579 (<u>https://github.com/BingqingCheng/ASAP</u>) are open-source.
- 580

# 581 ACKNOWLEDGMENTS

- 582 We thank Ian Ocampo for helpful discussions. We are grateful to reviewers for comments that improved
- the manuscript. We acknowledge the following grants: National Science Foundation (EAR-
- 584 1853388 to Lars Stixrude), National Natural Science Foundation of China (grant No. 22003050 to
- 585 Haiyang Niu), and the Research Fund of the State Key Laboratory of Solidification Processing
- 586 (NPU), China (grant No. 2020-QZ-03 to Haiyang Niu). This work used computational and storage
- 587 services associated with the Hoffman2 Shared Cluster provided by UCLA Institute for Digital
- 588 Research and Education's Research Technology Group.

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593 594	[1] C. Dorn A83 (2015).	n, A. Khan, K. Heng, J. A. D. Connolly, Y. Alibert, W. Benz, and P. Tackley, A&A 577,
595	[2] B. J. Fu	ulton and E. A. Petigura, The Astronomical Journal <b>156</b> , 264 (2018).
596	[3] C. Dorn	n and T. Lichtenberg, The Astrophysical Journal Letters 922, L4 (2021).
597	[4] L. Stixi	rude, Philos Trans A Math Phys Eng Sci 372, 20130076 (2014).
598	[5] R. G. K	Kraus, R. J. Hemley, S. J. Ali, J. L. Belof, L. X. Benedict, J. Bernier, D. Braun, R. E.
599	Cohen, G. W. C	Collins, F. Coppari et al., Science 375, 202 (2022).
600	[6] A. Bou	jibar, P. Driscoll, and Y. Fei, Journal of Geophysical Research: Planets 125,
601	e2019JE006124	4 (2020).
602	[7] S. Labr	cosse, J. W. Hernlund, and N. Coltice, Nature <b>450</b> , 866 (2007).
603	[8] A. Zerr	and R. Boehler, Science <b>262</b> , 553 (1993).
604	[9] D. L. H	Ieinz and R. Jeanloz, Journal of Geophysical Research: Solid Earth 92, 11437 (1987).
605	[10] Y. Fei,	C. T. Seagle, J. P. Townsend, C. A. McCoy, A. Boujibar, P. Driscoll, L. Shulenburger,
606	and M. D. Furn	ish, Nature Communications 12, 876 (2021).
607	[11] D. Frat	anduono, M. Millot, R. G. Kraus, D. K. Spaulding, G. W. Collins, P. M. Celliers, and J. H.
608	Eggert, Phys Ro	ev B 97, 214105 (2018).
609	[12] K. Um	emoto, R. M. Wentzcovitch, S. Wu, M. Ji, CZ. Wang, and KM. Ho, Earth and Planetary
610	Science Letters	<b>478</b> , 40 (2017).
612	[13] I. Lay,	J. Herniund, E. J. Garnero, and M. S. Thorne, Science <b>314</b> , 1272 (2006).
612	[14] SH. S	mm, Annual Review of Earth and Planetary Sciences <b>30</b> , 309 (2008).
617	and D I Prest	on Phys Pay Lett <b>0</b> / 105701 (2005)
615	[16] 7 $I$ $I$	in C P. Zhang X W Sun I B Hu T Song and V D Chu Physica Scrints 83 045602
616	(2011)	u, CK. Zhang, AW. Sun, JD. Hu, T. Song, and TD. Chu, Thysica Scripta 05, 045002
617	[17] C. Di P	Paola and J. P. Brodholt, Scientific Reports <b>6</b> , 29830 (2016).
618	[18] D. Alfè	M. J. Gillan, and G. D. Price. The Journal of Chemical Physics <b>116</b> , 6170 (2002).
619	[19] D. Alfè	e. Phys Rev Lett <b>94</b> , 235701 (2005).
620	[20] L. Voč	adlo, D. Alfè, G. D. Price, and M. J. Gillan, The Journal of Chemical Physics <b>120</b> , 2872
621	(2004).	
622	[21] J. Behle	er and M. Parrinello, Phys Rev Lett <b>98</b> , 146401 (2007).
623	[22] L. Zhai	ng, J. Han, H. Wang, R. Car, and W. E, Phys Rev Lett <b>120</b> , 143001 (2018).
624	[23] M. Yar	ng, T. Karmakar, and M. Parrinello, Phys Rev Lett 127, 080603 (2021).
625	[24] P. M. P	Piaggi and M. Parrinello, Phys Rev Lett 122, 050601 (2019).
626	[25] H. Niu,	, L. Bonati, P. M. Piaggi, and M. Parrinello, Nature Communications 11, 2654 (2020).
627	[26] G. A. T	ribello, M. Bonomi, D. Branduardi, C. Camilloni, and G. Bussi, Computer Physics
628	Communication	ns <b>185</b> , 604 (2014).
629	[27] S. Plim	pton, Journal of Computational Physics 117, 1 (1995).
630	[28] H. War	ng, L. Zhang, J. Han, and W. E, Computer Physics Communications 228, 178 (2018).
631	[29] H. Niu,	, P. M. Piaggi, M. Invernizzi, and M. Parrinello, Proceedings of the National Academy of
632	Sciences 115, 5	1348 (2018).
633	[30] H. Niu,	, Y. I. Yang, and M. Parrinello, Phys Rev Lett 122, 245501 (2019).
634	[31] P. M. F	Taggi and M. Parrinello, The Journal of Chemical Physics <b>150</b> , 244119 (2019).
035 626	[32] J. Deng	g and L. Suxrude, Geophysical Research Letters <b>562</b> , 116873 (2021).
030	$\begin{bmatrix} 33 \end{bmatrix}$ L. Stixi	rude and B. Karki, Science <b>310</b> , 297 (2005).
629	[34] A. F. B [35] C. Imb	alton, A. Analli, D. Giofrá, S. Klaos, I. Bahlar, and M. Cariotti, The Journal of Chamical
630	$\begin{bmatrix} 3 & 3 \end{bmatrix}  \textbf{U} \cdot \textbf{IIID}$	aizano, A. Anemi, D. Gione, S. Kiees, J. Beiner, and M. Cerioui, The Journal of Chemical (1730 (2018)
023	1 Hysics 140, 24	<i>113</i> 0 (2010).

- 640 [36] B. Cheng, R.-R. Griffiths, S. Wengert, C. Kunkel, T. Stenczel, B. Zhu, V. L. Deringer, N.
- 641 Bernstein, J. T. Margraf, K. Reuter *et al.*, Accounts of Chemical Research 53, 1981 (2020).
- 642 [37] H. Wang, L. Zhang, J. Han, and E. Weinan, Comput. Phys. Commun. 228, 178 (2018).
- 643 [38] S. Lorenz, A. Groß, and M. Scheffler, Chemical Physics Letters 395, 210 (2004).
- 644 [39] K. Hornik, M. Stinchcombe, and H. White, Neural Networks 2, 359 (1989).
- 645 [40] J. P. Perdew, A. Ruzsinszky, G. I. Csonka, O. A. Vydrov, G. E. Scuseria, L. A. Constantin, X.
- 646 Zhou, and K. Burke, Phys Rev Lett 100, 136406 (2008).
- 647 [41] G. Kresse and J. Furthmüller, Computational Materials Science 6, 15 (1996).
- 648 [42] G. Kresse and D. Joubert, Phys Rev B 59, 1758 (1999).
- 649 [43] E. Holmström and L. Stixrude, Phys Rev Lett 114, 117202 (2015).
- 650 [44] R. Scipioni, L. Stixrude, and M. P. Desjarlais, Proceedings of the National Academy of Sciences
  651 114, 9009 (2017).
- [45] J. Deng and L. Stixrude, Earth and Planetary Science Letters 562, 116873 (2021).
- [46] Y. Yoshimoto, Journal of the Physical Society of Japan 79, 034602 (2010).
- 654 [47] W. G. Hoover, Phys Rev A **31**, 1695 (1985).
- 655 [48] N. D. Mermin, Physical Review 137, A1441 (1965).
- 656 [49] A. Togo and I. Tanaka, Scripta Materialia 108, 1 (2015).
- [50] S. Baroni, S. de Gironcoli, A. Dal Corso, and P. Giannozzi, Rev Mod Phys 73, 515 (2001).
- 658 [51] E. Ito and T. Katsura, in *High Pressure Research: Application to Earth and Planetary*
- 659 *Sciences*1992), pp. 315.
- 660 [52] G. Shen and P. Lazor, Journal of Geophysical Research: Solid Earth 100, 17699 (1995).
- 661 [53] J. A. Akins, S.-N. Luo, P. D. Asimow, and T. J. Ahrens, Geophysical Research Letters 31 (2004).
- E. Knittle and R. Jeanloz, Geophysical Research Letters 16, 421 (1989).
- 663 [55] S. Tateno, K. Hirose, N. Sata, and Y. Ohishi, Earth and Planetary Science Letters **277**, 130 (2009).
- 665 [56] K. Hirose, R. Sinmyo, N. Sata, and Y. Ohishi, Geophysical Research Letters 33 (2006).
- 666 [57] R. A. Lange and I. S. E. Carmichael, Geochimica et Cosmochimica Acta 51, 2931 (1987).
- 667 [58] J. F. Stebbins, I. S. E. Carmichael, and L. K. Moret, Contributions to Mineralogy and Petrology
- **668 86**, 131 (1984).
- 669 [59] D. A. Young, *Phase diagrams of the elements* (University of California Press, Berkeley, 1991).
- 670 [60] J. Deng and K. K. M. Lee, Am Mineral 104, 1189 (2019).
- 671 [61] G. H. Wolf and R. Jeanloz, Journal of Geophysical Research: Solid Earth 89, 7821 (1984).
- 672 [62] S. M. Stishov, I. N. Makarenko, V. A. Ivanov, and A. M. Nikolaenko, Physics Letters A 45, 18673 (1973).
- [63] L. Stixrude, N. de Koker, N. Sun, M. Mookherjee, and B. B. Karki, Earth and Planetary Science
  Letters 278, 226 (2009).
- 676 [64] L. Stixrude and C. Lithgow-Bertelloni, Geophys J Int 162, 610 (2005).
- 677 [65] L. Stixrude and C. Lithgow-Bertelloni, Geophys J Int 184, 1180 (2011).
- 678 [66] L. Stixrude and C. Lithgow-Bertelloni, Geophys J Int 228, 1119 (2021).
- 679 [67] S. Stackhouse, L. Stixrude, and B. B. Karki, Earth and Planetary Science Letters 427, 11 (2015).
- 680 [68] L. Cobden, C. Thomas, and J. Trampert, in *The Earth's Heterogeneous Mantle: A Geophysical*,
- 681 *Geodynamical, and Geochemical Perspective*, edited by A. Khan, and F. Deschamps (Springer
- 682 International Publishing, Cham, 2015), pp. 391.
- 683 [69] C. Thomas, E. J. Garnero, and T. Lay, Journal of Geophysical Research: Solid Earth 109 (2004).
- F. Nimmo, in *Treatise on Geophysics (Second Edition)*, edited by G. Schubert (Elsevier, Oxford,
  2015), pp. 27.
- 686 [71] B. Grocholski, K. Catalli, S.-H. Shim, and V. Prakapenka, Proceedings of the National Academy 687 of Sciences **109**, 2275 (2012).
- 688